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# Sources of wintertime $PM_{2.5}$ at a major city in an Alpine Valley: The role of atmospheric dispersion and inversion dynamics

Mauro Masiol<sup>a,\*</sup>, Gianni Formenton<sup>b</sup>, Flavia Visin<sup>a</sup>, Alessandro Bonetto<sup>a</sup>, Manuela Rovea<sup>b</sup>, Silvia Ficotto<sup>b</sup>, Elisa Danesin<sup>a</sup>, Tommaso Toffanin<sup>a</sup>, Anita Maggiulli<sup>a</sup>, Maria Battistel<sup>a</sup>, Giovanna Mazzi<sup>a</sup>, Matteo Feltracco<sup>a</sup>, Andrea Gambaro<sup>a</sup>, Philip K. Hopke<sup>a,c,d</sup>

<sup>a</sup> Dept. Environmental Sciences, Informatics and Statistics, Università Ca' Foscari Venezia, Campus Scientifico, Via Torino 155, 30172, Mestre, VE, Italy

<sup>b</sup> ARPAV - Regional Agency for Environmental Protection and Prevention of Veneto, Via Lissa 6, 30174, Mestre, VE, Italy

<sup>c</sup> Dept. Public Health Sciences, University of Rochester Medical Center, Rochester, NY, 14642, USA

<sup>d</sup> Institute for a Sustainable Environment, Clarkson University, Potsdam, NY, 13699, USA

### HIGHLIGHTS

#### G R A P H I C A L A B S T R A C T

- PM<sub>2.5</sub> sources identified through PMF in a major city in an Alpine valley.
- Biomass burning account for 52% of PM<sub>2.5</sub>, secondary aerosol 21%, biogenic 20%.
- Persistent thermal inversion events identified from high-low altitude weather data.
- Biomass burning and secondary aerosol were higher in all inversion periods.
- Major effort needed to mitigate the biomass burning pollution in mountain valleys.

### ARTICLE INFO

Keywords: Particulate matter Chemical speciation Source apportionment Biomass burning Alpine environment



# ABSTRACT

Urban areas in mountain environments are generally located on valley floors surrounded by slopes, where mountain orography drives peculiar meteorology and atmospheric circulation. Also, persistent inversion dynamics may occur strongly affecting air pollution. This study characterizes the PM<sub>2.5</sub> pollution in a major city located in an Alpine valley (Belluno, Northeastern Italy) during the cold season (Autumn-Winter). Major particulate species (elemental and organic carbon, major inorganic ions) and minor/trace elements conventionally used as tracers for source apportionment were analyzed, including oxalate and specific PM<sub>2.5</sub>-bound tracers for biomass burning (K<sup>+</sup>, levoglucosan, mannosan, galactosan) and for primary biogenic organic aerosol (arabitol, mannitol, glucose). PM<sub>2.5</sub> sources are identified through positive matrix factorization and a series of post-processing tools. Results indicate that biomass burning, mostly emitted by residential wood combustion for domestic heating, is the major source (52% PM<sub>2.5</sub> mass concentration), followed by secondary aerosol (21%), biogenic aerosol (20%), traffic (4%), and dust resuspension (3%). The source contributions are assessed by accounting for the local meteorology. Insights into the dispersion or buildup of PM<sub>2.5</sub> sources were then

\* Corresponding author.

E-mail address: mauro.masiol@unive.it (M. Masiol).

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investigated by dispersion normalization. In addition, the possible effects of persistent thermal inversion events were evaluated by assessing the inversion strength from temperature profiles measured at multiple weather stations at different elevations with respect to the source contributions. Beyond the identification of those emission sources requiring further mitigation actions, this study also analyses the potential effects of local meteorology on  $PM_{2.5}$  pollution. The methodologies applied in this study can be easily adopted to other mountain environments for successful management of air quality.

#### 1. Introduction

Emissions of fine ambient particulate matter (PM25) dropped in the EU-27 Member States by 28% between 2005 and 2021 (EEA, 2023a) following increasingly stringent policies mandated by the European Union (EU) and the implementation of effective emission control technologies for several major emission source sectors, mostly energy generation, industry, and transport (Turnock et al., 2016). Despite widespread improvements in air quality across Europe (EEA, 2020), particulate pollution still exceeds some EU air quality standards (AQS) (e.g., Directive, 2008/50/EC; EU, 2008) and WHO guidelines (WHO, 2006; 2021) in some hotspots (EEA, 2020; 2023b). The non-attainment of AQS is expected to worsen further by 2030 since more stringent limit values are currently under discussion (EC, 2022) for a better alignment with the new WHO global air quality guidelines (WHO, 2021). The annual average PM<sub>2.5</sub> limit value will significantly drop from 20 to 10  $\mu g~m^{-3}$ , a limit comparable with US (9  $\mu g~m^{-3}$ ) or UK (10  $\mu g~m^{-3}$ ). Also, maximum daily concentrations for PM<sub>2.5</sub> will be introduced: 25  $\mu g~m^{-3}$ not to be exceeded more than 18 times per calendar year.

These new limit values are challenging to meet. For example, using Bayesian geostatistical regression models, Beloconi and Vounatsou (2023) showed that 47.5% of the total European population lived in areas where the revised EU thresholds were not attained during 2021. Given this issue, the main goal for research and local environmental protection agencies is to identify and quantify the major sources of PM mass. This urgent task is required to successfully plan future emissions reduction actions in time to meet the new AQS.

Northern Italy, particularly the Po Valley, is a known major European hotspot for PM (e.g., EEA, 2020; Masiol et al., 2017; Pivato et al., 2023; Ricciardelli et al., 2017; Scotto et al., 2021). In Veneto, a region encompassing the northeastern part of the Po Valley, numerous source apportionment studies have been conducted in the past decades to identify and quantify the most important sources of PM10, PM2.5 and PM<sub>1</sub> at several sites (e.g., Gregoris et al., 2021; Masiol et al., 2012, 2014; Squizzato et al., 2016, 2017; Stortini et al., 2009). However, most of the studies were performed in the surroundings of Venice, the regional capital city. A recent multiple-site study was conducted in 6 major cities across the region: 4 located in the main valley, one in the hilly area facing the valley and one in a mountain valley within the Alpine chain (Masiol et al., 2020). Results indicated common source profiles in all cities except in the mountain city. This difference was attributed to the potentially varying emission source profiles characterizing the mountain areas and differences in local meteorological conditions.

Urban areas in mountain environments are generally located on valley floors surrounded by slopes and are therefore "sheltered" from synoptic-wind systems. Thus, mountain orography drives peculiar meteorological and atmospheric circulation conditions (Whiteman, 2000) that is mostly dependent on local thermal forcing, such as local thermal winds (Cantelli et al., 2017; Giovannini et al., 2014; Gohm et al., 2009). Also, persistent thermal inversions may occur, especially in winter, trapping the pollutants close to the surface (Chazette et al., 2005; Largeron and Staquet, 2016). Additionally, natural and anthropogenic emission sources may differ from the other major cities in the Po Valley. For instance, natural biogenic emissions are expected to be significantly higher due to extended forests and Alpine grasslands as well as to changes in the ratios between anthropized vs. agricultural/natural landscapes. Furthermore, biomass burning (BB) in mountain areas is expected to account for a larger fraction of emissions from domestic heating because of the reduced availability of natural gas (smaller valleys are not connected), the availability of raw material (wood), and the greater propensity of the population to use traditional stoves.

This study characterized the  $PM_{2.5}$  pollution in a major city located in an Alpine valley, Belluno (BL), during the cold season (Autumn-Winter), identifying and quantifying the main potential sources. This source apportionment takes advantage of using specific organic tracers for BB and primary biogenic organic aerosol along with major particulate compounds and minor/trace elements conventionally used as tracers. Beyond the indication of which emission sources require further mitigation actions in a typical major Alpine city, this study also analyzes the potential effects of local meteorology on  $PM_{2.5}$  pollution. Insight into the dispersion or buildup of pollution sources were investigated by dispersion normalization (Dai et al., 2020) and by identifying periods with persistent thermal inversions.

#### 2. Materials and methods

#### 2.1. Study area

Valbelluna is a large mountain valley extending for ~50 km from NE to SW between the southern Alpine front ranges (pre-Alps) and the Dolomites. The height of surrounding mountains varies from 700 to 2500 m above sea level. It is one of the most populated areas of the Southeastern Alps, hosting around 120,000 inhabitants in two main cities, Belluno at the NE corner of the Valley (altitude 390 m, approx. 35,000 inhabitants over the whole municipality) and Feltre (20,000) at the SW corner, and many smaller towns scattered throughout the valley (Figs. S1a and b). The valley floor is heavily anthropized (Fig. S1c) with (i) major roads, (ii) a railway that is mostly electrified but also uses diesel-powered trains, (iii) small industries, mostly manufacturing home appliances, evewear, bathroom fixtures, and diary, and (iv) agriculture dominated by cultivated crops and orchards. Although the city is not a main tourist destination, it attracts artistic-cultural tourism and is indirectly linked to Alpine tourism both in winter (skiing) and summer. Beyond local emission sources, atmospheric transport of polluted air masses from neighboring anthropized areas such as the Po Valley represents an additional input of pollution over Alpine environments (Diemoz et al., 2019a; 2019b).

#### 2.2. Analytical

PM<sub>2.5</sub> was collected at an urban background station (BL, 46°8'30N, 12°13'3E, 394 m a.s.l.) managed by the Regional Environmental Protection Agency (ARPAV). The site is located within an urban green park in the city center (Fig. S1d) representative of city-wide air quality. PM<sub>2.5</sub> is continuously measured since 2009; in this study, the chemical speciation analysis is limited to October 1, 2020–February 14, 2021. A total of 133 daily samples were collected and weighted according to the UNI EN 12341:2014 standard using quartz fiber filters (Whatman QMA). PM<sub>2.5</sub> mass was measured gravimetrically (sensitivity ≤0.1 µg) with filter blanks and sampled filters conditioned for 48 h at 20 ± 1 °C and 50 ± 5% RH. To prevent positive or negative artifacts for organics and nitrate, photochemical reactions and biological processes, the sampler was kept under controlled temperature and filters were stored in clean Petri slides and refrigerated in the dark until analysis.

A 1 cm<sup>2</sup> punch was analyzed for EC and OC using a carbon aerosol analyzer (Sunset Lab, USA) adopting the EUSAAR2 protocol (Brown et al., 2017; Cavalli et al., 2010) and UNI EN 16909:2017. Two or three 1.5 cm<sup>2</sup> punches were extracted in ultrapure water (specific resistivity 18.2 M $\Omega$  cm) for 10 min under controlled temperature (20 °C) and then filtered through 0.22 µm PTFE filter disks. Aliquots of water-dissolved samples were analyzed for (i) major cations (Na<sup>+</sup>, NH<sub>4</sub><sup>+</sup>, K<sup>+</sup>, Mg<sup>2+</sup>,  $Ca^{2+}$ ) by ion-exchange chromatography equipped with conductivity detector; (ii) major anions (F<sup>-</sup>, Cl<sup>-</sup>, NO<sub>3</sub><sup>-</sup>, SO<sub>4</sub><sup>2-</sup>, oxalate (OXL)) by high-pressure ion chromatography equipped with conductivity detector and suppressor; and (iii) selected saccharides, levoglucosan (LVG), mannosan (MAN), galactosan (GAL), arabitol (ARB), mannitol (MNT), and glucose (GLC) by high-pressure ion chromatography equipped with a pulsed amperometric detector. Details are reported in SI section S1. The remaining portions of filters underwent elemental analysis (Al, V, Mn, Fe, Ni, Cu, Zn, As, Cd, Sn, Sb, Pb) by acid digestion (HNO3-HF-H2O2) in a microwave oven followed by ICP-OES and ICP-MS analyses (5300 DV and NexION 350X, PerkinElmer, USA, respectively). Details are provided in SI section S1.

In addition, PM<sub>10</sub> samples were collected alongside PM<sub>2.5</sub> according to the EN12341:2014 on OFF (Whatman OMA) and analyzed for particulate mass (UNI EN 12341:2014) and 8 polycyclic aromatic hydrocarbons (benz(a)anthracene (BaA), chrysene (Chry), benzo(b) fluoranthene (BbF), benzo(k)fluoranthene (BkF), benzo(a)pyrene (BaP), indeno(1,2,3-c,d)pyrene (IP), dibenzo(a,h)anthracene (DBahA) and benzo(g,h,i)perylene (BghiP)) following the UNI EN 15549:2008 and UNI CEN/TS 16645:2014 by HPLC interfaced with a multi-wavelength fluorescence detector (2695 series Alliance and 2475, Waters, USA; Masiol et al., 2013). PM<sub>10</sub>-bound PAH analyses were performed on selected filters and by extracting 2 to 3 filters together. Therefore, the data represents the average concentration over several days. Since PM-bound PAHs are mainly distributed into the submicrometric PM fraction (e.g., Lammel et al., 2011; Zhang et al., 2020), as confirmed by local studies in the mountain areas of Veneto (ARPAV, 2017), it is assumed that the  $PM_{10}$ -bound PAHs concentrations can be matched with the chemical-speciated PM2.5 data.

The quality of the analytical procedures was checked by blank controls, by evaluating detection limits (DLs), recoveries through reference materials, accuracy, and repeatability. Details in SI Section S1.

#### 2.3. Ancillary variables

NO, NO<sub>2</sub>, NO<sub>x</sub> (EN 14211:2012), and O<sub>3</sub> (EN 14625:2012) were also measured on an hourly basis at the same station. Weather variables including air temperature (°C), relative humidity (%), pressure (mbar), solar irradiation (W m<sup>-2</sup>) wind speed (m s<sup>-1</sup>) and direction were measured at a nearby station (AE) approx. 3.5 km NE of the city center. The water mixing ratio (g kg<sup>-1</sup>) was computed from the weather data.

The EC-tracer method (Cabada et al., 2004) was preliminarily applied to estimate the primary (POC) and secondary (SOC) organic carbon from OC/EC data. Details are reported in SI section S2. Additional variables were computed from the analytical data as proxies for specific atmospheric and/or aerosol properties; details are provided in SI section 3. Briefly, since hydroxyl radical is not measured, the oxidative potential of the atmosphere is indirectly assessed through the sum of measured atmospheric oxidants (OX=NO<sub>2</sub>+O<sub>3</sub> [ppb], Kley et al., 1999). The nitrogen oxidation ratio (NOR) depicts the degree of atmospheric conversion of NO<sub>2</sub> to particulate-phase nitrate (Colbeck and Harrison, 1984). The neutralization ratio (NR, Colbeck and Harrison, 1984; Bencs et al., 2008; Engelhart et al., 2011) describes the degree of neutralization of sulfate and nitrate by ammonium.

The CALMET model was used to estimate the planetary boundary height (Hmix) and atmospheric stability. Details are reported in SI section S4. The estimation of the weather-related impacts on the net energy demand for heating was assessed by computing the heating degree-days (HDD; Thom, 1954; Quayle and Diaz, 1980; Spinoni et al., 2018). Details in SI section S5 and Table S1. Essentially, HDDs are proportional to the amount of energy needed to heat the internal environment in a cold climate to a specified base temperature for a given period (15.5 °C and daily basis in this case; Spinoni et al., 2015). Consequently, the HDD value reflects the potential biomass consumption for domestic heating each day.

#### 2.4. Mass closure

Mass closure was assessed by summing the concentrations of the analyzed chemical species with assumptions to account for elements not directly measured, mostly hydrogen and oxygen in both organic and inorganic compounds. Existing methods are reviewed by Chow et al. (2015). The compositions of major chemical components (organic matter (OM), elemental carbon (EC), sulfate, nitrate, ammonium (SNA), crustal material (CRU), sea-salt aerosol (SSA), and other trace elements and ions (OTHERS)) were assessed (details in SI section S6).

#### 2.5. Source apportionment

 $PM_{2.5}$  sources were identified and quantified by USEPA PMF version 5 (Hopke, 2015; Paatero, 1997; Paatero and Tapper, 1994). Data below DLs were set as DL/2, with an uncertainty of 5/6 of the corresponding DL. Uncertainties for data > DLs were carefully assessed by propagating errors from all steps of the adopted analytical procedures, including analytical uncertainties in samples and blanks, recoveries, and sampler performances. Uncertainty for the total variable (PM<sub>2.5</sub>) was set to 300% of the concentration to prevent it from driving the model (Kim et al., 2003). The optimal solution was identified based on the best practice reported in the relevant literature (Belis et al., 2014; Brown et al., 2015; Hopke, 2016; Reff et al., 2007) and the evaluations from the built-in diagnostic tools. Rotational ambiguity was assessed through displacement (DISP); uncertainties were estimated by bootstrap (BS, replicates = 200, Min. correlation, r = 0.6), and the combined BS-DISP (Paatero et al., 2014) was calculated to check the overall stability of the solutions.

The remaining data analyses were computed by R 4.3.3 (R Core Team, 2022). The location of potential local sources was assessed by mapping the average pollutant concentrations by wind speed and direction as a continuous surface in bivariate polar plots (Carslaw and Ropkins, 2012).

## 2.6. Dispersion normalization

Dispersion normalization using the ventilation coefficient (VC) was used to account for the effects of meteorological variations on air quality and to improve source apportionment results (e.g., Chen et al., 2022; Dai et al., 2020; Sofowote et al., 2021). Essentially, the atmospheric dispersion is assumed to be proportional to the product of wind speed (ws) and boundary layer height:

$$VC_i = ws_i \times Hmix_i$$
 (1)

where i denotes each measurement period (days). Since ws and Hmix are provided on an hourly basis, the daily averages are used. The daily concentration values are then normalized:

$$C_{VC,i} = C_i \times \frac{VC_i}{\overline{VC}}$$
(2)

where  $C_{VC,i}$  is the normalized concentration,  $C_i$  is the measured concentration for the period i, and  $\overline{VC}$  is the average VC over the whole sampling period.

# 2.7. Detection of persistent inversion periods

Persistent thermal inversions negatively affect air pollution in Alpine valleys. Since vertical temperature profiles are not experimentally

measured, a simple estimation of the inversion strength across the valley was computed using air temperature from nearby weather stations at different elevations. This method was adapted from Largeron and Staquet (2016), who reported persistent inversion in the Grenoble area (French Alps) related to elevated  $PM_{10}$  concentrations. This method returns similar results to other methods available in literature such as the heat deficit evaluation (Whiteman et al., 1999).

Hourly weather data from 8 stations (Fig. S1; Table S2) managed by ARPAV were initially investigated. These ground-based stations are located at different elevations (264 m–1603 m ASL) within the Valbelluna or in the surrounding mountains. Since only three stations (FE, SG, AE) are located in the center of the valley floor at 264–379 m a.s.l., these stations were paired with three more stations located at high elevation (MA, MC, FA, at 1415–1603 m a.s.l., respectively). The horizontal homogeneity of air temperature was preliminarily checked by correlating the hourly-resolved weather data measured at low- and high-elevation stations during the study period. Assuming horizontal homogeneity of the air temperature field over the area, the bulk temperature gradient ( $\Delta T/\Delta H$ ) for each pair of stations was computed and used as a proxy of the stability of the boundary layer into the valley.  $\Delta T/\Delta H$  represents the ratio between the difference in air temperature to the difference in height between the high- and low-elevation stations:

$$(\Delta T/\Delta H)_{i,k} = \frac{T_{i,k(\text{ high})} - T_{i,k(\text{ low})}}{H_{k(\text{ high})} - H_{k(\text{ low})}}$$
(3)

where i denotes each measurement period (hourly data, in this case), k is a station pair [1,3],  $H_{k(high)}$  and  $H_{k(low)}$  represent the station elevation and  $T_{i,k(high)}$  and  $T_{i,k(low)}$  represent the air temperature measured at the time i of the pair k at the high and low station respectively. Table S3 summarizes the characteristics of station pairs.

 $\Delta T/\Delta H$  is considered a valid indicator of the stability of the boundary layer when an inversion occurs. According to Largeron and Staquet (2016),  $\Delta T/\Delta H$  >0 indicates that an inversion is present between the two stations but not necessarily that the entire layer between the two stations is filled by the inversion. Therefore, inversions were identified by evaluating the temporal patterns of daily-averaged  $\Delta T/\Delta H$  ratios. Periods with daily  $\Delta T/\Delta H$  >0 for more than 3 days were considered persistent temperature inversions.

#### 3. Results and discussion

Fig. 1 shows the distributions of concentrations of PM<sub>2.5</sub> and analyzed PM<sub>2.5</sub>-bound species; Fig. S2 reports the distributions of meteorological, gaseous pollutant, and other ancillary variables. The averaged concentration of major (>1% wt) analyzed species followed the order (in  $\mu$ g m<sup>-3</sup>: OC (6.2) > NO<sub>3</sub><sup>-</sup> (2.7) > EC (1.9) > LVG (0.9) > NH<sub>4</sub><sup>+</sup> (0.8) > SO<sub>4</sub><sup>2-</sup> (0.7) > K<sup>+</sup> (0.4) > Al (0.3). The carbonaceous fraction (OC + EC) accounted for 49% of the total PM<sub>2.5</sub> mass (8.1  $\mu$ g m<sup>-3</sup>). EC and OC were strongly correlated (r = 0.93), were generally highly correlated with anions, K<sup>+</sup>, ammonium, all organics (including BB tracers but not GLC), Zn and Cd but not well correlated with other cations and elements (Fig. S3).

The average OC/EC ratio was 3.2 (range 2–5) with the higher values recorded in January. An OC to EC ratio for the primary sources (OC/EC)<sub>p</sub> of 2.42 was found adopting the EC-tracer method (Fig. S4). POC (average 4.5  $\mu g$  m  $^{-3},$  28% of the total  $PM_{2.5}$  mass) largely exceeded SOC (1.7  $\mu$ g m<sup>-3</sup>, 9%), showing SOC/POC ratios ranging from 0 to 1.1 (average 0.36). Water-soluble inorganic ions (WSIIs) represented the second largest contribution to the total  $PM_{2.5}$  mass (26%, 4.6 µg m<sup>-3</sup>). The major WSII component was the sulfate-nitrate-ammonium (SNA) aerosol forming in the atmosphere following well-established thermodynamic processes involving the H<sub>2</sub>SO<sub>4</sub>-HNO<sub>3</sub>-NH<sub>3</sub> system (Seinfeld and Pandis, 2016). SNA accounted for approx. 22% of the total PM<sub>2.5</sub> mass (4  $\mu$ g m<sup>-3</sup>) and 77% of total WSIIs. Nitrate accounted for 16% of PM<sub>2.5</sub> mass indicating that ammonium nitrate is the dominant form of the secondary inorganic aerosol (SIA) as commonly reported in Northeastern Italy (Masiol et al., 2015). Generally, all ions were moderately to strongly correlated to each other, except  $Mg^{2+}$  and  $Ca^{2+}$  (Fig. S3). They are also well correlated with organic tracers except GLC.

Among the common BB tracers (Bhattarai et al., 2019; Puxbaum et al., 2007; Schkolink et al., 2006; Simoneit, 2002), LVG exhibited higher average concentrations followed by K<sup>+</sup>, MAN, and GAL (Fig. 1). Very high correlations were found among the BB tracers (Figures S3 and S5a-c), with coefficients of determination ranging from 0.94 (K<sup>+</sup>-LVG) to 0.98 (LVG-MAN, MAN-GAL). LVG also exhibited a strong linear relationship ( $r^2 = 0.88$ ) with POC but a much lower ( $r^2 = 0.58$ ) with SOC (Figs. S5d and e), clearly reflecting the primary nature of this specific marker.

GLC, MNT and ARB are often used as tracers for primary biogenic organic aerosol (Marynowski and Simoneit, 2022; Pashynska et al., 2002; Samaké et al., 2019; Yttri et al., 2007). Specifically, GLC is associated with plant materials (e.g., pollen, fragments of leaves) or soil emissions, while MNT and ARB are generally associated with airborne fungal spores (e.g., Rathnayake et al., 2017; Samaké et al., 2019; Wan et al., 2019). Strong positive and significant relationships (Figs. S3 and S6) were found between MNT and ARB (r = 0.83) and between polyols (defined as MNT + ARB) and OC (r = 0.88), while no correlation was found between GLC and polyols. The concentrations of GLC, MNT and ARB are in the range of those collected at different locations and seasons, as reviewed by Golly et al. (2019).

OXL, a derivative form of oxalic acid, is generally associated with secondary formation processes, representing an end-product of the oxidation of organic gas-phase species (Ervens et al., 2004). The concentration varied between 28 and 253 ng m<sup>-3</sup> (average 96 ng m<sup>-3</sup>), similar to concentrations measured at an alpine remote site in France (95.5 ng m<sup>-3</sup> during all seasons; Golly et al., 2019) but lower than those measured in urban areas in Northern Italy, (Milan 122 ng m<sup>-3</sup> in winter; Perrone et al., 2012). OXL accounted for 0.7–4.9% of total OC (average 1.8%) and was strongly positively correlated with SNA ions (0.81 < r < 0.88; Fig. S3), stressing its main secondary origin. However, it is not strongly correlated with POC (r = 0.44) or SOC (r = 0.55) from the EC-tracer method. Ratios between specific chemical species are often used as proxies for inferring sources and processes and are discussed in Section S7 and Figs. S7 and S8.



Fig. 1. Boxplots of the analyzed compounds and elements in  $PM_{2.5}$  at BL. For each species, the boxplots show: line = median, box = inter-quartile range, whiskers =  $\pm 1.5^{\pm}$  inter-quartile range; red cross = arithmetic means; outliers and extremes are not shown.

## 3.1. Nitrate aerosol formation

Insights into the nitrate aerosol formation arise from some simple steps and a few assumptions. Ammonium sulfate is preferentially and irreversibly formed when sulfuric acid and ammonia are available, taking priority over ammonium nitrate because of its low vapor pressure. Under ammonium-rich conditions ( $NH_4^+/SO_4^{2-}$  molar ratio >2), it can be assumed that ammonium sulfate is the preferred form of sulfate, otherwise ammonium sulfate and bisulfate may form in ammoniumpoor conditions, increasing the aerosol acidity. The reconstruction of the ammonium concentration from nitrate and sulfate indicated a quite good fit ( $r^2 > 0.95$ ; Fig. S9a), with slopes close to unity but positive intercepts. However, ammonium-deficient conditions were generally observed (points over the 1:1 line). This ammonium imbalance was also responsible for the neutralization ratios below unity (Fig. S2), as previously reported in Belluno (Masiol et al., 2015). Despite this, the cation-to-anion ratio in charge equivalents exhibits a perfect ion balance (slope 0.996,  $r^2 = 0.97$ ; Fig. S9b) even if Na<sup>+</sup> was not included (below DL). This result is possibly explained by the formation of mixed salts.

As a simple way to better elucidate the factors affecting ammonium nitrate formation, the linear regression between the  $NH_4^+/SO_4^{2-}$  and  $NO_3^-/SO_4^{2-}$  molar ratios for ammonium-rich conditions (r<sup>2</sup> = 0.69; Fig. S9c) returns an intercept of 1.14 and a slope of 0.79, which ratio (1.44) below 2 further supports the assumption that part of the sulfate existed in other forms, such as ammonium bisulfate (Pathak et al., 2009) or mixed salts. After the sulfate aerosol formed, ammonium nitrate forms when more ammonium is available, i.e., when the excess- $NH_{4}^{+}$  > 0 (details in SI section S3). Fig. 2 shows that the excess- $NH_4^+$  ranges from -4 to 233 nmol m<sup>-3</sup> (average 32 nmol m<sup>-3</sup>) with 77% of samples under ammonium-rich regimes. Despite the uncertainty of the adopted method (which does not consider the gas-phase acids and ammonia and analytical uncertainties), possible sampling artifacts (limited by the low air temperatures and strict sampling protocols) and the large variability of data, the nitrate to excess-NH<sub>4</sub><sup>+</sup> ratios are mostly over the 1:1 line, again indicating the possible formation of mixed nitrate salts.

#### 3.2. Mass closure

In this study, the primary (POM) and secondary (SOM) organic matter were separately estimated from the POC and SOC from the ECtracer method, considering the different oxidation between the primary and secondary organic aerosol. A 1.2 factor was applied to POC,



**Fig. 2.** Scatter plots of nitrate vs. excess-ammonium (in molar concentrations). The plot also shows other variables computed in the study (details in SI section S2): the vertical red line shows the split between ammonium-poor and -rich conditions; the dimension of points is proportional to the nitrogen oxidation ratio (NOR); the color of points shows the neutralization ratio (NR). (For interpretation of the references to color in this figure legend, the reader is referred to the web version of this article.)

assuming the composition of fresh aerosol in urban areas (Chow et al., 2002), while SOC was multiplied by 1.8, a value indicative of aerosol from BB emissions after ageing for  $\sim$ 3 days from smog chamber studies (Lim et al., 2019). These values agree with the 1.6 value used in previous studies in BL for total OC (Masiol et al., 2020). Crustal material and salt were computed according to equations in Hand (2011). Secondary inorganic aerosol was simply assessed as SNA.

Results (Fig. 3a) show good mass reconstruction (slope = 0.96) and fit ( $r^2$  = 0.94, RMSE 2.2 µg m<sup>-3</sup>) with concentrations of the various PM<sub>2.5</sub> chemical components over the whole sampling period following the order (in µg m<sup>-3</sup>): POM (5.5) > SNA (4.2), crustal material (3.4) > SOM (3.1) > EC (1.9) > other species (0.5) > sea-salt (0.1) (Fig. 3b). The relative contributions of the major chemical components were quite constant during the sampling period (Fig. S10), but the absolute concentrations of POM, SOM, SIA, EC, and other species significantly increased from October to January (Fig. 3c), as indicated by the Kruskal-Wallis test (p < 0.05) and its pairwise comparisons using Dunn's posthoc test (p < 0.0002). No other differences were statistically significant.

#### 3.3. Source apportionment

Variables were screened for their information content (Paatero and Hopke, 2003): a total of 26 variables were inputted as "strong", while 3 variables with signal-to-noise  $\leq 1$  (Al, Cu, As) were set as "weak" thereby tripling their uncertainties. Four days were excluded from the analysis: 12 October because significantly worsened the diagnostics and exhibited high-scaled residuals for multiple species, probably due to construction activities or contamination, and 5-7 December because concentrations of PM25 and multiple species were below DLs due to intense precipitation. Solutions were explored from 3 to 7 factors. The most physically plausible results were obtained with a 5-factor solution, for which the minimum Q values were evaluated over 200 runs. The selected solution also exhibited the  $Q_{true}$  value closest to the theoretical ( $Q_{exp}$ ) value  $(Q_{true}/Q_{exp} \approx 0.97)$  and optimal diagnostics: (i) all runs converged; (ii) all species had stable Q values over 200 runs; (iii) the absolute scaled residuals were symmetrically distributed and the number of values exceeding  $\pm 3$  was minimized; (iv) modeled total variable (PM<sub>2.5</sub>) was successfully predicted ( $r^2 = 0.97$ ); (v) no swaps by factor were found in DISP with % dQ = 0; and (vi) swaps in BS were minimized (<10%) with no unmapped factors. The G-space showed no edges (Paatero et al., 2005), thus FPEAK solutions were not investigated. No constraints (Brown et al., 2015) were applied after extensive exploration showed no improvement in the profiles. According to previous literature (Zíková et al., 2016), the min-max ranges calculated by DISP under the constraint that the difference (dQ = base-modified) is no greater than the  $dQ_{max} = 4$  were used to assess the boundaries of source profiles.

Results were interpreted based on the presence of known tracers in the source profiles (Fig. 4), the ratios of selected chemical species in the profiles (Fig. 5), the local directionality of factors shown by polar plot analysis (Fig. S11), the temporal patterns of contributions throughout the sampling period (Fig. S12), and the correlations with PAHs, weather and ancillary variables (Fig. S13).

#### 3.3.1. Secondary aerosol

Factor 1 presents a high share of ammonium (74% of specie sum), nitrate (46%), sulfate (37%) and OC (11%), but also contains significant fractions of OXL and As. It strongly correlates (r > 0.65) with PM<sub>10</sub> and moderately correlates (0.45 < r < 0.65) with most PAH congeners (Fig. S13). This source represents secondary aerosols mostly composed of ammonium nitrate, ammonium sulfate/bisulfate and organics. The nitrate-sulfate-ammonium association remains unchanged by increasing the number of factors. Generally, the chemistry of secondary nitrate competes with secondary sulfate, but the presence of a cumulative source is not surprising in this study. Secondary nitrate is highly favored by cold temperatures (period average: 4.3 °C) and high relative humidity (period average: 90%) (Seinfeld and Pandis, 2016) usually



**Fig. 3.** Results of the mass closure: a) linear regression showing the fit between the measured and reconstructed  $PM_{2.5}$  masses; b) the distributions of the 7 major  $PM_{2.5}$  components (details in SI section S6); c) cumulative concentrations of the 7 major  $PM_{2.5}$  components reconstructed by the mass closure (sorted by increasing average contributions over the whole sampling period). (For interpretation of the references to color in this figure legend, the reader is referred to the web version of this article.)



Fig. 4. PMF factor profiles. Left y-axis: bars represent mass contribution of the base run; open circles represent the mean DISP values with the error bars providing the range (minima and maxima values) of DISP values. Right y-axis: red filled squares show factor contributions in percent of species sum. (For interpretation of the references to color in this figure legend, the reader is referred to the web version of this article.)

recorded in the study area in winter and transition seasons.

Conversely, sulfate production, despite being more thermodynamically stable, is enhanced by photochemistry and is mostly dominating during warm seasons (Squizzato et al., 2013). The presence of a relatively large fraction of OC reflects the production of secondary organic aerosol (SOA) enhanced in the presence of pre-existing acidic particles (Deng et al., 2021; Hallquist et al., 2009; Kuwata et al., 2015; Shrivastava et al., 2017), resulting in internally mixed aerosols. This factor has the highest OC/EC ratio (Fig. 5), approximately 3-folds higher than the primary OC/EC computed with the EC-tracer method. This result



**Fig. 5.** PMF results: a) boxplots of the source contributions; b) stacked bar plot showing the average source contributions of the 5 sources by months and over the whole sampling period; c) relative contribution of the 5 sources by months and over the whole sampling period; d), e), and f) ratios among chemical species in the PMF profiles (the grey bars refer to the concentration in the base run; error bars show the uncertainty of the ratios as 5-95th percentile ranges over the 200 BS runs). (For interpretation of the references to color in this figure legend, the reader is referred to the web version of this article.)

indicates the uptake of SOA on primary particles. The 22% of total OXL in this factor stresses its secondary nature (Zhou et al., 2015). OXL is principally generated in the atmosphere via atmospheric oxidation of higher homologues of long-chain diacids as well as from other pollution-derived organic precursors in gas, heterogeneous, and aqueous phases (Kawamura and Bikkina, 2016; Yang et al., 2022) and was found internally mixed with aged BB particles (Zhang et al., 2017). The relatively high shares of As, Sn and Fe are unusual for a secondary factor. However, the factor does not present swaps with any other factors and the association does not split until a 7-factor solution was investigated, returning poorer diagnostics. Thus, it is unlikely that this chemical association represents distinct sources due to a poorly resolved solution.

The secondary aerosol source accounts for 21% (DISP min-max range 19–29%) of total PM<sub>2.5</sub> mass (Fig. 5). Polar plot analysis does not show a clear directionality but high concentrations by moderate winds blowing from ENE and S (Fig. S11). The time series shows no clear patterns but the alternation of periods with low and high concentrations. During 21–23 October, a severe SA peak occurred with concentrations up to 32  $\mu$ g m<sup>-3</sup> (Fig. S12), i.e., 7-fold higher than the remaining study period.

#### 3.3.2. Biomass burning

Factor 2 is responsible for more than half (52%) of total PM<sub>2.5</sub> mass (Fig. 5) and shows very high shares of traditional BB tracers: anhydrosugars (75-78%), K<sup>+</sup> (73%) as well as high contributions of OC (65%) and EC (62%). It was interpreted as BB emissions mostly composed of mixing of organics, soot and potassium, according to the literature (Lee et al., 2016). The factor concentration constantly increases from October to January (Fig. 5) following HDD (r = 0.69) and inversely to air temperature (r = -0.69) (Fig. S13). Although this source represents combustion sources likely dominated by residential wood combustion (RWC) (Vicente and Alves, 2018), the possible contribution of open crop residue fires cannot be neglected in autumn. Crop residue burning is not allowed but is still utilized in the post-harvest season. RWC is a known major aerosol source in Europe (Cincinelli et al., 2019; Kukkonen et al., 2020; Yttri et al., 2019). The BB factor is also strongly associated with NO<sub>x</sub> and PAHs (0.68 < r < 0.92), stressing its primary origin from combustion. A stepwise multiple linear regression was

performed between PMF sources (independent variables) and the sum of 8 analyzed congeners (dependent variable). BB was the only independent variable statistically significant (p < 0.05). Adopting the Akaike information criterion as an estimator of prediction errors (in both directions), the best model was found with only BB, indicating that almost all PAHs are emitted by this source.

BB does not present clear directionality (Fig. S11) but showed increased concentrations during low wind regimes and moderate easterly winds, compatible with a local origin. The OC/EC ratio of the source profile (BS range 3.3–3.9) is slightly higher than the primary OC/EC ratio found with the EC-tracer method but generally agreed with data reported in literature for emissions of woodstoves burning wooden logs (mainly pine) (2.8  $\pm$  1.3; Zhang et al., 2013) or spruce wood logs (1.2–11.5; Orasche et al., 2012), and pellet stoves burning wood pellet (0.9–4.2; Vicente et al., 2015).

The BB factor also exhibits a moderate share of OXL (28%). Sources of oxalate include primary and secondary emissions from biogenic and anthropogenic sources, such as BB (Kundu et al., 2010). For example, Xu et al. (2020a) showed similar results (BB contributed 25% of OXL during a typical biomass burning period using single particle source analysis). Beyond the primary emission, a possible ageing of BB emissions cannot be excluded from this factor. Recent literature reports that organic aerosol from BB may undergo fast chemical ageing even under winter-time periods with low photochemical activity or nighttime (e.g., Kodros et al., 2020; Yazdani et al., 2023). In addition, the high relative humidity (period average 90% RH, with 34% of sampling days with daily average > 95% RH) may favor the formation of secondary organic aerosols in the aqueous phase (Gilardoni et al., 2016).

The BB factor is also associated with high fractions of fluoride (76%), chloride (43%), and nitrate (48%), which are known to be emitted by BB processes (Calvo et al., 2015; Fine et al., 2001; Jayarathne et al., 2014; Song and Liu, 2023; Vicente and Alves, 2018). Also, it shows high fractions of ARB (59%) and MNT (49%). Despite being typical markers of primary biological aerosols (Marynowski and Simoneit, 2022; Samaké et al., 2019; Xu et al., 2020b), these sugar alcohols are also found in association with primary BB emissions in wintertime (Xu et al., 2019). Moderate shares of Fe, Zn and Cu are not surprising since Schmidl et al. (2008) reported that these elements are emitted by wood stove

combustion of common woods grown in Alpine regions.

#### 3.3.3. Biogenic aerosol

The third factor accounts for 20% of PM2.5 mass and is mostly composed of OC. It also exhibits high fractions of EC, GLC,  $Ca^{2+}$ ,  $Mg^{2+}$ ARB, MNT, and OXL. Although LVG, K<sup>+</sup> and nitrate show moderate base run contributions, their wide DISP intervals toward null contribution indicate that these species have limited associations with this factor. Instead, the presence of GLC, MNT and ARB suggests biogenic emissions. This interpretation is also supported by the decreasing monthly contributions from October to January. The factor also shows a moderate share of elemental carbon but the DISP interval takes EC down to  $\sim 66$ ng  $m^{-3}$ , which is quite a low concentration. Also, the biogenic factor is likely an indicator of SOA from biogenic volatile organic compounds (BVOCs). These BVOCs need a place to condense, and elemental carbon particles will serve as a preferred surface on which to condense because of their unique chain aggregate shape (high surface to volume). Possible secondary origin is also supported by the high shares of OXL and sulfate, indicating possible atmospheric aging. This way, the EC share may be explained by mixing with other material during resuspension or coemission along with BB source in stoves during cold-start phases. Polar plots analysis (Fig. S11) does not show clear directionality but higher concentrations during moderate to high wind days, which is compatible with the resuspension of material of biological origin from the surrounding forests and green areas.

## 3.3.4. Traffic and resuspension dust

The last two factors account for a small fraction of  $PM_{2.5}$  mass (4% and 3%, respectively; Fig. 5) showing constant monthly contributions and similar profiles. They exhibit significant fractions of Al, transition, and post-transition metals (V, Mn, Fe, Ni, Cu, Zn, Sn and Pb). On the contrary, OC, EC, ammonium, potassium, BB tracers and GLC show wide DISP ranges and are not likely associated with these factors. Metals can be emitted from multiple sources, such as exhaust-related sources (engine corrosion, fuel and lubricant combustion, catalytic converters, particulate filters) as well as from non-exhaust sources (tire and brake wear, muffler ablation) (Amato, 2018; Fussell et al., 2022; Pant and Harrison, 2013). The main differences between the two sources are the presence of high Sn and Sb concentrations in factor 4 and high shares of Ca<sup>2+</sup> and Mg<sup>2+</sup> in factor 5. Considering the DISP ranges, nitrate has a moderate value in factor 4, while chloride and fluoride are in factor 5.

Insights upon their interpretation come from wind analysis. Factor 4 shows no directionality and higher concentrations for low-moderate winds, suggesting a source widespread over the city. Traffic is the most likely source. This interpretation is also supported by the similarity with the factor profile for traffic reported in major cities in Veneto (Masiol et al., 2020) considering the specific elements. Factor 5 increased during high wind regimes (correlation with wind speed r = 0.40) suggesting dust resuspension. This interpretation is supported by the higher shares of crustal elements ( $Ca^{2+}$ ,  $Mg^{2+}$ ) linked to calcareous soils, dolomite and limestones present in the study area, and shares of chlorine, compatible with road salts used as anti-icing and the higher share of Zn, generally associated with tire wear.

# 3.4. The role of specific sources on the exceedance of European standards

The average  $PM_{2.5}$  concentration was 17  $\mu g~m^{-3}$  with 22 days exceeding a daily concentration of 25  $\mu g~m^{-3}$ . The PMF source contributions when the forthcoming daily limit is met vs. exceedances show statistically significant differences only for SA and BB (nonparametric Wilcoxon signed rank test) at p < 0.01 (Fig. S14a). Average increases of 7.6  $\mu g~m^{-3}$  of SA and 9.9  $\mu g~m^{-3}$  of BB were recorded during exceedances with respect to days with daily limit value met (Figs. S14b and c).

Thus, these sources must be the main target of air pollution mitigation strategies. By contrast, TRA and RD are not significantly different possibly because current strategies for the mitigation of air pollution (traffic restrictions based upon EURO vehicle's emissions standard) are in force. Thus, specific air pollution mitigation policies for  $PM_{2.5}$  should be mostly implemented or strengthened to meet newly proposed limit values during the cold season, targeting the emissions from domestic appliances burning wood.

Wilcoxon sign-rank tests indicated that the PM<sub>2.5</sub> mass concentration and the contributions of secondary aerosol and BB were statistically significantly (p < 0.01) higher in days with ammonium-rich with respect to ammonium-poor conditions. No significant differences were found for other sources. In addition, all days with PM<sub>2.5</sub>>25  $\mu$ g m<sup>-3</sup> occurred during ammonium-rich conditions. Thus, the ammonia availability plays a key role over the nitrate aerosol formation and aerosol mass. Since EIs for 2019 in Valbelluna municipalities (INEMAR, 2022) estimate that 90% of ammonia (762 Mg y<sup>-1</sup>) is emitted by agriculture, of which 90% (688 Mg y<sup>-1</sup>) come from manure management regarding nitrogen compounds. Therefore, it is necessary that further actions are also implemented at a regional and supra-regional level to reduce ammonia emissions for minimizing the secondary nitrate aerosol.

# 3.5. The effects of atmospheric dispersion

Generally, the dispersion normalization concentrations of  $PM_{2.5}$  were lower than un-normalized data (Fig. S15a), mostly due to the inverse relationship between Hmix (decreasing from October to February) and  $PM_{2.5}$  concentration (increasing). A period of high wind regimes (11–14 February) led to very high VC-normalized values at the end of the sampling period. Another high-wind period occurred on 5–7 December along with strong precipitation, but these samples were removed from PMF due to concentrations below DLs.  $PM_{2.5}$  concentrations accounting for the VC show a different monthly pattern (Fig. 6 and S15b,c), with higher  $PM_{2.5}$  concentrations in October and February (only 14 days are available in February). These differences were statistically significant (nonparametric Wilcoxon signed rank test) during all periods, except in February (Fig. 6).

Examining the single sources, BB shows the main differences, with significantly lower VC-normalized concentrations in November–January. Also, BIO and RD showed statistically significant differences between raw and VC-normalized contributions with higher concentrations in October and lower in November–December. This result stresses again the relationship of BIO and RD, being generated by the resuspension of natural and anthropogenic materials, respectively. TRA shows lower VCnormalized concentrations in November and December, while SA only in January. In summary, the dispersion normalization results show that the BB source is more affected by the VC factor because the ventilation conditions of the valley favor the increase in concentrations (normalized VC lower than those detected experimentally). The two resuspension sources are also influenced by atmospheric dispersion, but the October concentrations are effectively dispersed, probably because the higher average daily height of the planetary boundary layer.

#### 3.6. The effects of persistent inversions

A preliminary evaluation of the air temperature shows a strong correlation (0.68 < r < 0.99) among the 8 selected weather stations (Fig. S16) with the higher values found between the 3 low-elevation (FE, SG, AE, TO; 0.96  $\leq$  r  $\leq$  0.99) and the 3 high-elevation (MA, MC, FA; 0.94  $\leq$  r  $\leq$  0.99) stations. The linear relationships of air temperatures confirm this strong correlation (Figs. S17a–c). By contrast, the linear relationships between pairs of high-elevation and low-elevation stations (Figs. S17d–e) show temperature differences between the valley and the mountaintop are much less correlated and exhibit a clear edge that closely follows the dry adiabatic lapse rate (DALR, 9.8 °C km<sup>-1</sup>). These latter scatterplots also show that most of the observations lie between the DALR and the 1:1 isothermal line. However, a limited number of points lie below the isothermal, i.e., the high elevation station recorded higher temperature values than the low elevation.



Fig. 6. PM<sub>2.5</sub> and PMF source contributions: measured PM<sub>2.5</sub> and contribution modelized by PMF are in grey; dispersion-normalized concentration in pink (For interpretation of the references to color in this figure legend, the reader is referred to the web version of this article.).

The time series at hourly and daily resolutions (Fig. S18) also show generally lower temperatures in the high-elevation stations, but also periods with potential thermal inversions. Similar results were reported by Largeron and Staquet (2016) in the Grenoble Valley, France, emphasizing that local processes do not strongly account for the spatiotemporal variability of air temperature inside the valley and that synoptic scale fluctuations may play a major role, especially close to the summits. Thus, horizontal gradients in air temperature are negligible with respect to the vertical changes. This approximation allows considering air temperature horizontally homogeneous across the Valley, being the low-elevation stations AE and FE spaced by 32 km and the high-elevation stations FA and VP spaced by 43 km.

Based on the horizontal homogeneity assumption for the air temperature at the low- and high-elevation stations, 3 pairs of weather stations were selected to be representative of the  $\Delta T/\Delta H$  ratios over all the valley (FA-AE for the NE valley, MC-SG for the center, and MA-FE for SW; characteristics of station pairs are summarized in Table S3) and the temperature gradient ( $\Delta T/\Delta H$ ) for each pair of stations was computed. The  $\Delta T/\Delta H$  at different station pairs were strongly correlated (Fig. S19) between pairs of stations, indicating that potential inversions filled the whole valley.

Fig. 7a and b shows the hourly and daily  $\Delta T/\Delta H,$  respectively. Six



**Fig. 7.** Identification of the persistent thermal inversion periods: a) hourly-resolved  $\Delta T/\Delta H$ , b) daily averaged  $\Delta T/\Delta H$ , and c) concentrations of PM2.5 measured at Belluno (raw data in black, VC-normalized in red) and in Feltre. (For interpretation of the references to color in this figure legend, the reader is referred to the web version of this article.)



Fig. 8. Differences between the average (points) min and max (bars) PM<sub>2.5</sub> and PMF source contributions measured during the persistent inversion periods and the concentrations measured 2 days before. Raw data are shown in black; VC-normalized concentrations in grey.

periods with daily  $\Delta T/\Delta H >0$  for more than 3 days were considered persistent temperature inversions (Table S4). Generally, these periods were characterized by or just followed high atmospheric pressure days (Fig. S12), whose effect was compatible with the instauration of inversion conditions in valleys. The PM<sub>2.5</sub> concentration and the PMF source contributions during the 6 inversion periods are then compared with the concentrations measured two days before. Results indicate a general increase in air pollution during the inversion (Fig. 8). PM<sub>2.5</sub> concentrations were higher in all inversion periods (0–20 µg m<sup>-3</sup>) as well as SA (0–8 µg m<sup>-3</sup>) and BB (0–10 µg m<sup>-3</sup>). The remaining sources did not show clear patterns. Fig. 8 also shows the VC-normalized concentrations, whose effects on air pollution were less evident, indicating that the dispersion normalization is also able to account for the effects of thermal inversions in valleys.

# 4. Conclusions

This study identifies the main sources of  $PM_{2.5}$  in a major city located within an Alpine valley and quantifies the potential effects of atmospheric dispersion and persistent temperature dynamics. The results show that biomass combustion (52%) and the secondary aerosol fraction (21%) mostly contribute to the  $PM_{2.5}$  mass concentration. These sources are responsible for excessive  $PM_{2.5}$  concentrations, i.e., they contribute in an even more significant way when considering periods of heavy pollution, here identified as days with  $PM_{2.5}$  concentrations higher than 25 µg m<sup>-3</sup>, which corresponds to the new daily AQS proposed to the European Commission. Other sources include biogenic materials (20%), traffic (4%), and resuspension dust (3%).

Although limited by the uncertainty of the mixing layer height modeling, complex landscape and wind dynamics in mountain environments, the calculation of concentrations normalized by the ventilation coefficient suggests that biomass burning is significantly affected by local atmospheric dispersion. Also, biomass burning and secondary aerosol increase significantly during periods in which a thermal inversion is persistently established in the valley, conditions that likely occur in winter. This result is particularly worrying since mountain valleys are frequently prone to long-lasting inversion conditions, mostly occurring overnight and during cold seasons, or in periods with weak synoptic forcing.

Under this view, considering the effects of local meteorology, the new standard for air quality planned for 2030 will be difficult to achieve without strong intervention and the implementation of successful mitigation strategies at a local scale aiming to successfully reduce the local emissions from space heating, mostly from wood burning. Thus, to attain the coming EU standards, a major effort will be needed to switch home heating to cleaner fuels or improve the combustion efficiency of wood stoves and boilers. These recommendations are not inconsistent with climate policies encouraging biomass as a climate-neutral (nonfossil) fuel but suggest revising and integrating the EU policies to mitigate the potential adverse effects on air quality. In addition, emissions of secondary aerosol precursors must also be lowered but on an interregional scale encompassing the Alpine area. In particular, nonattainments occur during ammonium-rich conditions. Thus, particular care should be given to the control of ammonia emissions.

Beyond the local importance of these results for attaining forthcoming European AQ standards, this study demonstrated that the combination of a traditional source apportionment approach with simple data analysis on weather parameters is useful to quantitatively account for the effects of atmospheric dispersion potential and persistent thermal inversions in an Alpine valley.

#### CRediT authorship contribution statement

Mauro Masiol: Writing - review & editing, Writing - original draft, Validation, Methodology, Investigation, Funding acquisition, Formal analysis, Data curation, Conceptualization. Gianni Formenton: Writing - review & editing, Writing - original draft, Validation, Resources, Methodology, Investigation, Formal analysis, Conceptualization. Flavia Visin: Validation, Methodology, Formal analysis. Alessandro Bonetto: Validation, Methodology, Formal analysis. Manuela Rovea: Validation, Methodology, Formal analysis. Silvia Ficotto: Validation, Methodology, Formal analysis. Elisa Danesin: Validation, Methodology, Formal analysis. Tommaso Toffanin: Validation, Methodology, Formal analysis. Anita Maggiulli: Validation, Methodology, Formal analysis. Maria Battistel: Writing - review & editing, Validation, Methodology, Formal analysis. Giovanna Mazzi: Writing - review & editing, Validation, Methodology, Formal analysis. Matteo Feltracco: Writing - review & editing, Validation, Methodology, Formal analysis. Andrea Gambaro: Writing - review & editing, Project administration, Funding acquisition, Conceptualization. Philip K. Hopke: Writing - review & editing, Writing - original draft, Methodology, Investigation, Conceptualization.

#### Declaration of competing interest

The authors declare that they have no known competing financial interests or personal relationships that could have appeared to influence the work reported in this paper.

#### Data availability

Data will be made available on request.

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## Appendix A. Supplementary data

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